



# *Chieseiceras dolomiticum* n. sp. (Ammonoidea) and its significance for the calibration of the Triassic platform interior stratigraphy at Latemar (Southern Alps, Italy)

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## Abstract

The species *Chieseiceras dolomiticum* n. sp. along with other ammonoid species (*Latemarites latemarensis*, *Halilucites rusticus*) helps constraining the age of the cyclic platform interior portion at Latemar (Dolomites, northern Italy). The comparison of ammonoids from the Latemar platform with the fossil record in basinal successions in the Southern Alps (including the Ladinian GSSP section at Bagolino) and Hungary suggests that the entire rhythmically bedded Latemar interval is entirely late Anisian in age. More than 400 stratigraphical metres of bedded shallow water carbonates at Latemar are found to correspond to less than 5 m of siliceous nodular limestone in the pelagic succession at Bagolino. The refined correlation will be significant for the reassessment of the cyclic Latemar stratigraphy in the light of new geochronological calibration of the South Alpine Middle Triassic.

**Keywords** Middle Triassic · Dolomites · Latemar · Carbonate platforms · Ammonoids

## Introduction

Middle Triassic carbonate platforms in the Southern Alps are unique for the rather unusual but relatively widespread occurrences of age diagnostic “pelagic” fossils (ammonoids and *Daonella*, thin-shelled pelecypods). These fossils provide the most important ties for a detailed correlation of platform carbonates with coeval basinal sedimentary rocks (Brack and Rieber 1993; Fantini Sestini 1994, 1996; De Zanche et al. 1995; Manfrin et al. 2005).

Among the South Alpine platforms (Fig. 1), Latemar is exceptional because numerous fossil-bearing horizons occur in a pile of spectacularly exposed, regularly bedded and shallow dipping platform interior strata. The setting represents a temporal succession and helps in calibration

because it offers the rare opportunity to assess the factors controlling the cyclic bedding in shallow water carbonates.

In the past three decades, the sedimentology at Latemar has been studied in detail (Goldhammer et al. 1987, 1990; Egenhoff et al. 1999; Preto et al. 2004; Zühlke 2004; Emmerich et al. 2005; Marangon et al. 2011). However, divergent opinions on bedding characteristics and their durations triggered a heated debate, in the literature sometimes referred to as the “Latemar controversy” (Hinnov and Goldhammer 1991; Brack et al. 1996; Preto et al. 2001; Mundil et al. 2003; Zühlke et al. 2003; Kent et al. 2004; Meyers 2008). The discussion centres around the interpretation of (sub) metre-scale Latemar cycles driven by sea-level oscillations at Milankovitch or sub-Milankovitch frequencies.

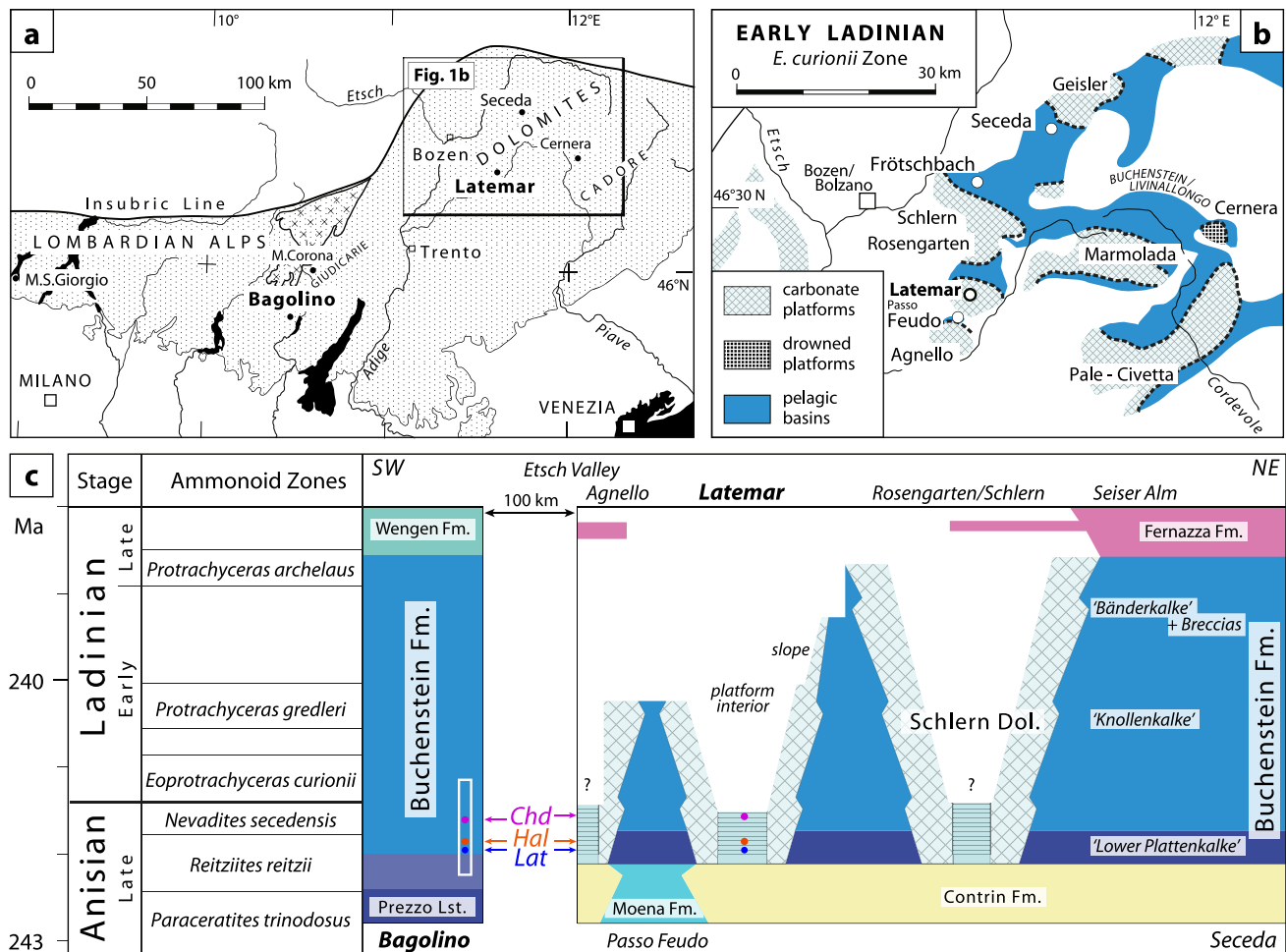
In basinal successions, the ammonoid record coeval with the Latemar interval is well known from the Buchenstein Formation and other formations across the Southern Alps (e.g., Brack and Rieber 1993; Mietto et al. 2003; Preto et al. 2005) as well as from similar lithological units in the Balaton Highland in Hungary (Vörös 1998; Pálffy et al. 2003). Much of the palaeontological information is known from the GSSP section of the Ladinian Stage at Bagolino (Brack et al. 2005). Stratigraphic data from well-studied coeval Hungarian sections can be projected in detail onto Bagolino (Vörös et al.

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**Fig. 1** **a** Overview of the central and western Southern Alps showing the location of the Bagolino reference section in eastern Lombardy as well as the map of **b**. **b** Distribution of carbonate platforms and basins in the central and western Dolomites during the late Anisian—early Ladinian. **c** Chronostratigraphic scheme for upper Anisian to Ladinian formations in eastern Lombardy (Bagolino) and in the northwestern

Dolomites (modified after Storck et al. 2019). The white bar in the Bagolino column marks the stratigraphic interval shown in Fig. 2. The stratigraphic positions of key ammonoids at Latemar and in the basinal successions are indicated (Lat: *Latemarites latemarensis*, Hal: *Halilucites rusticus*, Chd: *Chieseiceras dolomiticum*)

2008). This correlation adds further biostratigraphic constraints on the duration of the Latemar cycles.

In a first step towards the reevaluation of the bedding systematics at Latemar in the light of an improved recent geochronology-based age calibration of the South Alpine Middle Triassic (Wotzlaw et al. 2018; Storck et al. 2019), this article briefly summarises the most important biostratigraphic age information for the platform interior succession at Latemar. This is complemented by new data including an ammonoid species (*Chieseiceras dolomiticum* n. sp.) that helps constraining the temporal allocation of the platform interior succession at Latemar. The material, upon which the new species is based, was discovered two decades ago by students of the University of Heidelberg, on the summit of Cimon del Latemar (Dolomites, northern Italy). A rock slab packed with small fossils (predominantly ammonoids) and with the size of about two bricks was donated by Rainer

Zühlke (then at the University of Heidelberg) to Hans Rieber (Palaeontological Institute of the University of Zürich) for preparation and study. A joint visit of the locality allowed us to verify the bedrock origin of this material.

### The Latemar platform

Latemar was a possibly isolated carbonate platform with a diameter of around 3 km. The current exposures display a bedded core from which slope deposits dip radially away with up to 35° steep clinoforms. At the toe of the slopes, turbidites and breccias interfinger with basinal carbonates of the Buchenstein Formation. At Latemar, only parts of this facies change are preserved. The complete platform-basin transition is exposed in the neighbouring Rosengarten platform (e.g., Maurer 2000).

The initial shape of the Latemar platform may have resembled a horseshoe with an embayment towards the deeply incised Val Sorda on the southeastern side of Latemar (Preto et al. 2011; Gramigna et al. 2013). Along the steep flanks of this valley, the platform interior succession is fully exposed and accessible. The bedded platform interior exceeds 600 m in thickness and has been divided into six lithozones (Egenhoff et al. 1999). The coarsely stratified Lower Platform Facies (LPF) and early slope clinoforms follow on top of the (pre-Latemar) Contrin Formation and are overlain by dm- to m-scale bedded intervals that are distinguished on the basis of relative abundance of exposure surfaces and tepee structures. In ascending order these include the Lower Tepee Facies (LTF), Lower Cyclic Facies (LCF), Middle Tepee Facies (MTF), Upper Cyclic Facies (UCF) and Upper Tepee Facies (UTF). The succession from LCF to UTF is over 450 m thick and hosts the disputed cyclic bedding patterns.

## Ammonoids at Latemar

### Platform interior beds

Ammonoids from Latemar have long been known (e.g., Richthofen 1860; Mojsisovics 1882; Philipp 1904; Wilckens 1909, Koken 1911; Bubnoff 1921; Assereto 1969; Gaetani et al. 1981) but most of these findings were from allochthonous blocks or from platform slope portions with unclear stratigraphy. However, in the past 30 years, ammonoids have been discovered also in numerous places located in the stratigraphic succession of the platform interior beds (Brack and Rieber 1993; De Zanche et al. 1995; Mietto and Manfrin 1995; Mundil et al. 2003; Zühlke 2004, Preto et al. 2004; Manfrin et al. 2005, see Fig. 3 in their article for locations). The fossil localities in the shallow water carbonates are in most cases less than one kilometre away from coeval platform margins. The shells of “pelagic” organisms were likely imported from the open sea and concentrated by (possibly storm-induced) currents. In the regularly stratified cyclic facies ammonoids and other fossils locally form coquinas in tepee cavities and in intertepee depressions (Egenhoff et al. 1999; Manfrin et al. 2005). The stratigraphically oldest level (“Lastei di Valsorda” in Brack and Rieber 1993) is a peculiar accumulation lagerstätte which Preto et al. (2011) consider to be part of the clinostratified infilling of a closed or semi-closed fault-bounded depression (their “tectonic blue hole”) at the level of the LPF/LTF.

To date, ammonoids from over ten stratigraphically distinct levels constrain the biostratigraphic age of the cyclically bedded platform interior at Latemar to a short time span in the late Anisian. This corresponds to the upper *R. reitzi* to *N. secedensis* ammonoid zones of Brack et al.

(2005) and to the *Avisianum* to *Secedensis* (= *Serpianensis*) subzones of Manfrin et al. (2005).

Ammonoids from the oldest level including *Latemarites latemarensis* (= *L. bavaricus* of Manfrin et al. 2005), *Aplococeras avisianum*, *Parakellnerites rothpletzi*, *Hungarites zalaensis* and other taxa have been listed and illustrated in Brack and Rieber (1993). For comparable fossils and those from younger levels we refer to Manfrin et al. (2005, their localities TP 9–La 109; see also Preto et al. 2004 for the distribution of fossil layers). Our collection confirms the findings of these authors. Own finds include the following taxa:

- *Parakellnerites rothpletzi* and other forms of *Parakellnerites*, *Aplococeras misanii*, *Hungarites*, *Tropigastrites*, *Gymnites* from the Torre di Pisa (TP 9) locality;
- *Parakellnerites waageni*, *Parakellnerites* sp., *Aplococeras misanii*, ? *Stoppaniceras* from Cima di Valsorda (LCV 7-12, L 1, L 2);
- *Aplococeras misanii*, Celtitidae from a level corresponding to or somewhat above La 105.

*Chieseiceras dolomiticum* n. sp. described below is from the stratigraphically highest level of platform interior strata at Cimon del Latemar (ca. 1 m below its top layer). There, it occurs in association with numerous other ammonoids (*Praepinacoceras airaghii*, *Norites* cf. *dieneri*, *Aplococeras misanii*, *Parakellnerites* sp., *Megaphyllites* sp., *Epigymnites* sp., *Sturia* cf. *forojulensi*, Celtitidae) along with other cephalopods (*Michelinoceras*, *Mojsisovicstheutis*, ? *Syringonautilus*), gastropods and dasycladalean algae, in a laterally restricted accumulation of fossils.

From the stratigraphically largely equivalent level La 0 at Cimon del Latemar (i.e., 10 m below the top) Manfrin et al. (2005) also mention *Proarcestes*, *Paranevadites* and “*Detoniceras*”. It is important to note that both faunas from Cimon del Latemar lack truly Ladinian elements such as *Eoprotrachyceras*, *Protrachyceras* and *Arpadites*. A fauna similar to Cimon del Latemar is known also from fossil lenses in the Esino Limestone of Val Parina in the Bergamask Alps (Fantini Sestini 1996). Ammonoids from outcrop S1060B in Val Parina comprise *Praepinacoceras airaghii* along with specimens of *Parakellnerites*, *Norites*, *Aplococeras*, *Megaphyllites*, *Epigymnites* and other taxa. This and other nearby lenses in Val Parina are older than those with faunas from the *E. curionii* Zone.

### Platform slope portions

At Latemar ammonoids along with different forms of *Daonella* also occur in numerous places in platform slope deposits. However, due to the irregular platform geometry and complex architecture of the slope portions (Emmerich et al. 2005; Preto et al. 2011) fossils from these places can

be used only as secondary constraints for the age of the platform interior succession. Some of the faunas occur in sedimentary dykes and as fills of cavities (Brack and Rieber 1993; Emmerich et al. 2005; Marangon et al. 2011).

The most prolific fossil-bearing slope spot hitherto known is at Isugadoi (Koken 1911; Brack and Rieber 1993). *Daonella* and ammonoids including representatives of *Nevadites*, *Falsanolcites*, *Aplococeras misanii* are indicative for the *N. secedensis* Zone (Brack et al. 2005). Similar *Daonella* were found also west of Gamsstallscharte.

Ammonoids from Schenon (Latemar Ostgipfel) including *Arpadites* (Philipp 1904; Wilckens 1909) and *Protrachyceras* cf. *longobardicum* (= *P. ex gr. ladinum* of Gaetani et al. 1981) are younger (“*P. gredleri*” Zone and younger levels) and might have been part of fissure fills (Gaetani et al. 1981). The same holds true for *Daonella taramellii* found in a sedimentary dyke crossing the slope deposits on the crest ca 140 m east–northeast of the highest point of Schenon (= fauna L4 in Brack et al. 1996).

### Correlation of basinal records with the platform interior at Latemar

While many ammonoids reported from Latemar are long ranging forms or have insufficient diagnostic features allowing precise correlations, a few distinct Latemar ammonoid species occur within more restricted intervals in the basinal sediments (Fig. 2). Among these are representatives of *Latemarites* and *Halilucites* along with the new species of *Chieseiceras* described below. *Latemarites latemarensis* and *Chieseiceras dolomiticum* n. sp. occur in the lowest (at the level of Lower Tepee Facies, LTF, or somewhat below) and in the uppermost fossiliferous levels (Upper Tepee Facies, UTF), respectively, and are thus most important for constraining the base and the top of the cyclic platform interior portion at Latemar. From an intermediate position in this succession (Middle Tepee Facies, MTF; Cima di Valsorda—level LCV 12) *Halilucites rusticus* has been illustrated by Manfrin et al. (2005).

In basinal successions of the Southern Alps, *Latemarites* sp. has been mentioned from two sections in the eastern Dolomites (Preto et al. 2005). The basinal intervals there are more expanded and can be correlated only approximately with the reference sections of the Buchenstein Formation at Seceda or Bagolino. However, the species *Latemarites latemarensis* itself is well known from Hungarian localities (Vörös 1998; Pálffy et al. 2003). At Felsőörs it occurs between the horizons with *Reitziites* and *Ticinities* (Vörös et al. 2008), in an interval that is strongly expanded through interbedded, up to metre-thick volcanoclastic layers (Fig. 2). The *Reitziites* and *Ticinities* levels are known also from the GSSP section at Bagolino. There, only mm to cm-thin volcanoclastic layers are found in the

interval between *Reitziites* and *Ticinities* but the amount of carbonate layers compares well with the cumulative thickness of limestone beds and nodules at Felsőörs (see also Brack et al. 2005). *Halilucites rusticus* is known from Felsőörs in a stratigraphical position below the *Ticinities*-layer as at Bagolino (Fig. 3; Mietto et al. 2003; Brack et al. 2005). A well-preserved specimen of *Chieseiceras dolomiticum* n. sp. (Fig. 3) is from a bed immediately above the Tc-tuff layers at Bagolino and was hitherto referred to with open nomenclature (*Chieseiceras* sp.; Brack et al. 2005). This level is within the *N. secedensis* ammonoid Zone.

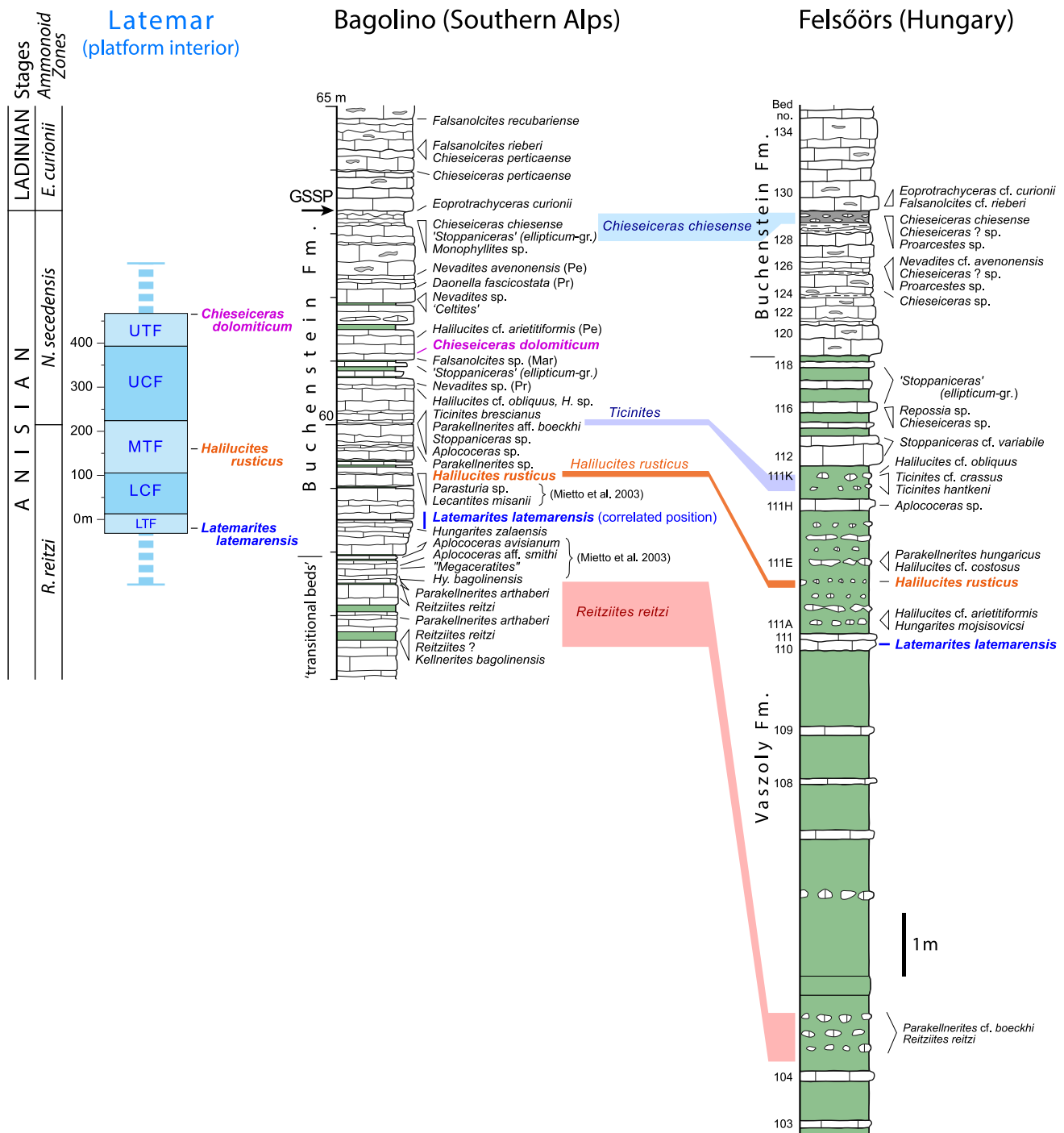
The sequence of the three marker species from Latemar thus correlates with the combined Bagolino–Felsőörs record and the proportional vertical distance between the levels of these ammonoids is comparable. Moreover, this finding is in agreement with the distribution of other ammonoid taxa as far as their ranges can be judged.

In the Buchenstein Formation of the Southern Alps ornamented ammonoids (trachyostraca) are often known to occur at single levels or in rather restricted intervals. Around the Anisian–Ladinian boundary, a possibly typical range for the appearance of a single species may be seen in the distribution of the most abundant taxon *Chieseiceras chiesense*. In the nodular limestone facies of the Buchenstein Fm. of eastern Lombardy and Giudicarie (including Bagolino), the vertical distribution of this species known from numerous sections does not exceed 50 cm (Brack and Rieber 1986, 1993).

Even when doubling this value for the ranges of marker ammonoid taxa used here for correlation relative to the Bagolino setting, the proposed platform-basin comparison implies a correspondence of the > 400 m thick cyclic platform interior succession at Latemar with less than five stratigraphic metres of the Bagolino column (Fig. 3). The time interval for the cyclic facies at Latemar may thus be substantially shorter than hitherto assumed. This is shorter than the recent correlation proposed by Spahn et al. (2013) and adopted also by Franceschi et al. (2016). The correlation of Spahn et al. (2013) is based on the identification of magnetic reversals combined with the tracing of volcanoclastic layers. However, at least nine ash layers have been identified to date in the platform interior succession at Latemar (Zühlke 2004; Preto et al. 2005) and the correlation to the numerous volcanoclastic layers in the basinal successions is not straightforward. Moreover, the existence of a reversed polarity interval at Latemar as supposed by Spahn et al. (2013) is questionable (Kent et al. 2004).

### Conclusions

The reevaluation of ammonoid data from Latemar and coeval basinal sections in the Southern Alps and Hungary confirms the short duration of the cyclic platform interior

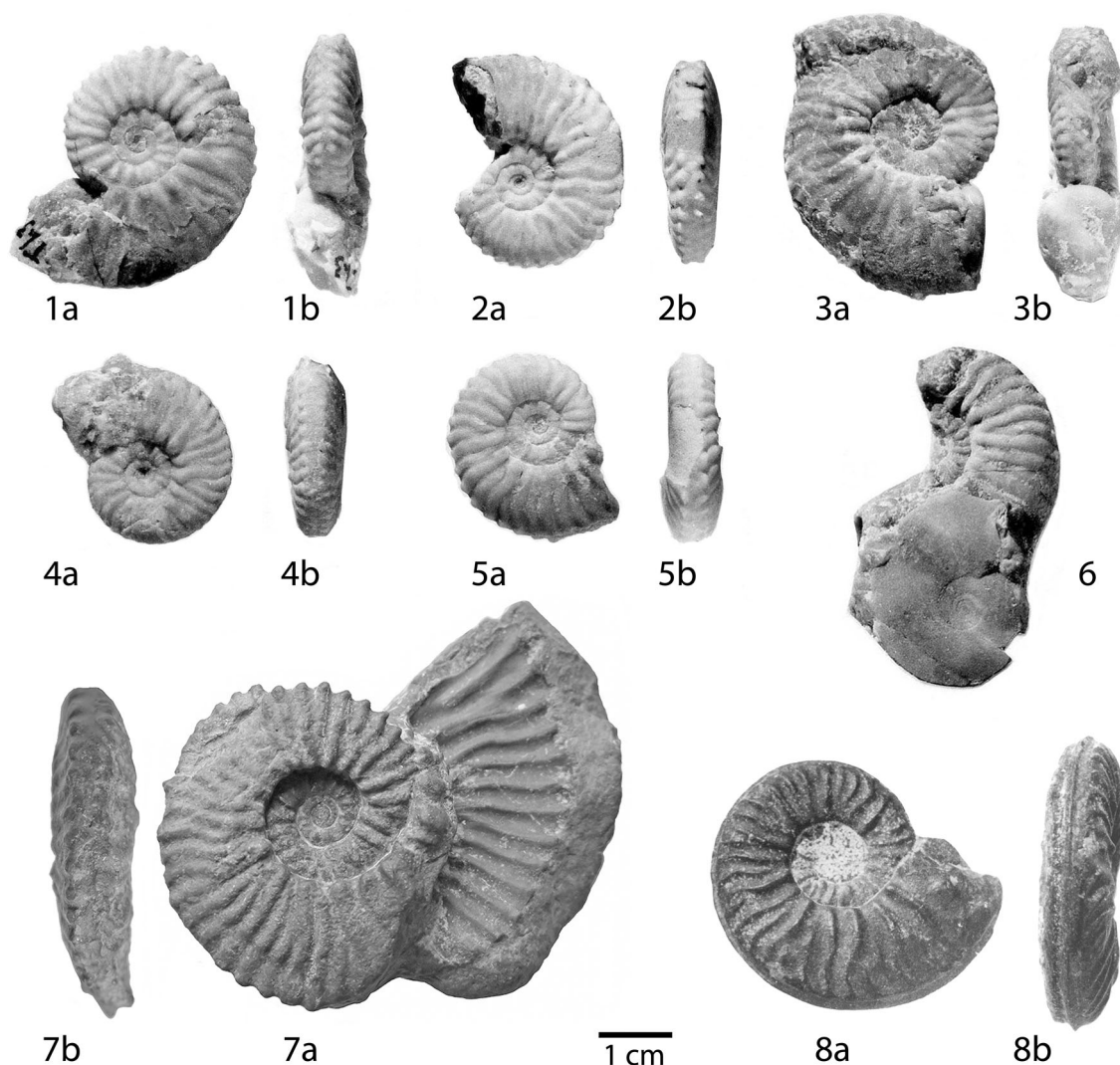


**Fig. 2** Correlation of the platform interior succession at Latemar (LTF Lower Tepee Facies, LCF Lower Cyclic Facies, MTF Middle Tepee Facies, UCF upper Cyclic Facies, UTF Upper Tepee Facies; metre scale as for section “Latemartürme” in Zühlke et al. 2003) and basal sections at Bagolino and Felsőörs (Hungary). Bagolino is updated and modified after Brack et al. (2005). A few important ammonoids and *Daonella* projected from nearby sections are labelled

interval at Latemar. The new ammonoid species *Chieseiceras dolomiticum* n. sp. adds to the reliability of this correlation and reduces the range of uncertainty

(Pe: Pèrtica, Pr: Prezzo, Mar: Marcheno). Felsőörs is redrawn after Vörös et al. (2008, Figs. 3, 4). See their article for details of the correlation also of deeper parts of the sections below the *Reitziites reitzi* level. Green colour highlights volcanoclastic layers. The supposed maximal range of the Latemar succession relative to the Bagolino reference section is indicated (stippled blue line) (colour figure online)

especially for the top of the Latemar succession as preserved today.



**Fig. 3** Ammonoids from the peak of Cimon del Latemar and from the Buchenstein Fm. at Bagolino. 1–6: *Chieseiceras dolomiticum* n. sp., uppermost Latemar Lst. (top of Upper Tepee Facies) at Cimon del Latemar, 1: PIMUZ 25701 (holotype); 2: PIMUZ 36535; 3: PIMUZ 36536; 4: PIMUZ 36537; 5: PIMUZ 36538; 6: *Ch. dolomiticum* with

*Praepinacoceras airaghii* Fantini Sestini, 1996, PIMUZ 36539. 7: *Chieseiceras dolomiticum* n. sp., Buchenstein Fm., Bagolino (PIMUZ 36534). 8: *Halilucites rusticus*, Buchenstein Fm., Bagolino (PIMUZ 36533). The specimen collected in 2001 corresponds to the in situ imprint illustrated in Mietto et al. (2003, pl. 1, Fig. 17)

The resulting correlation implies that the > 400 m thick cyclic platform interior interval at Latemar is Late Anisian in age and unlikely exceeds the top of this stage. This interval rather corresponds to less than five stratigraphic metres in the GSSP-reference section at Bagolino (Fig. 2). The proposed refined correlation indicates that the Latemar succession is shorter than hitherto assumed. This correspondence will be relevant for a future calibration and interpretation of the bedding patterns at Latemar in terms of absolute ages and time span.

### Systematic description

Subclass Ammonoidea Zittel, 1884

Order Ceratitida Hyatt, 1884

Superfamily Ceratitoidea Mojsisovics, 1879

Family Ceratitidae Mojsisovics, 1879

Subfamily Nevaditinae Tozer, 1994

Genus *Chieseiceras* Brack and Rieber, 1986

*Chieseiceras dolomiticum* n. sp.

Figure 3, 1–6.

*Synonymy*

? 1998 *Chieseiceras chiesense* (Mojsisovics)—Vörös, A. tab. 8, Fig. 4a + b

v 2003 *Chieseiceras* sp.—Bechstädt et al. (2003), Field-Trip to Latemar, fig III-15/1.

**Table 1** Morphological features of specimens of different species of *Chieseiceras*. Multipliers indicate observations on quarter whorls

Specimen	(A) Diameter (cm) at location of measurements	(B) Umbilical width (%)	(C) Umbilical nodes per half whorl	(D) Marginal nodes per half whorl	Ratio (d/c)
<i>Ch. dolomiticum</i> (PIMUZ 25701, Fig. 3/1) holotype, Latemar	2.6	38	9	18	2.0
<i>Ch. dolomiticum</i> (PIMUZ 36535, Fig. 3/2) Latemar	2.9	33	8.5	20	2.4
<i>Ch. dolomiticum</i> (PIMUZ 36537, Fig. 3/4) Latemar	2.35	32	9	20.5	2.3
<i>Ch. dolomiticum</i> (PIMUZ 36538, Fig. 3/5) Latemar	2.3	35	8	17.5	2.2
<i>Ch. dolomiticum</i> (PIMUZ 36534, Fig. 3/7) Bagolino	4.4	36	10 <sup>+</sup>	21	2.1
<i>Ch. chiesense</i> (B&R 1986, pl.1/Fig. 7) Bagolino	4.3	32.5	2 × 6 12	2 × 15 30	2.5
<i>Ch. chiesense</i> (B&R 1986, pl.1/Fig. 1) Bagolino	~ 3.5	31.5	2 × 4.5 9	2 × 11 22	2.4
<i>Ch. perticaense</i> holotype Pèrtica Bassa (B&R 1986, pl.3/Fig. 1+4)	~ 5.0	—	15	31	2.1
<i>Ch. perticaense</i> (B&R 1986, pl.3/Fig. 3) Pèrtica Bassa	2.9	31	11	25	2.3
<i>Ch. pemphix</i> (Mojsisovics 1982, pl.39/Fig. 9) San Salvatore	3.7	40	8	20	2.5

v 2005 *Chieseiceras* sp.—Brack et al., Fig. 7 (Bagolino column)

**Holotype** Specimen PIMUZ 25701, (Fig. 3/1a, b; the ammonoid is figured also on Fig. 2 of Mundil et al. 2003, but there with a wrong PIMUZ number).

**Name** Referring to the locality of discovery in the Dolomites (northern Italy).

**Type locality and horizon** The slab from which the holotype and many other small ammonoids and a few additional fossils were isolated was situated stratigraphically about 1 m below the top of the succession a few metres southeast of the cross on the summit of Cimon del Latemar (WGS84 coord.: 46°22'52.0"N, 11°34'30.2"E). The laterally restricted accumulation of fossils was found in light-coloured calcareous sediments of the highest part of the

Upper Tepee Facies (UTF) in the Latemar Limestone (Sciliar Fm. in Gramigna et al. 2013).

**Material** Six small specimens and several fragments isolated from a slab of Latemar Limestone (platform facies) from Cimon del Latemar and two specimens from the pelagic Buchenstein Formation at Bagolino (eastern Lombardy, northern Italy).

Specimens labelled with PIMUZ-numbers are stored in the collection of the Palaeontological Institute and Museum of the University of Zurich.

**Diagnosis** Evolute *Chieseiceras* with tabulate whorl section and broad, rectiradiate ribs. The ribs start at the umbilical border with a more or less blunt node. In different positions of the inner half of the flank secondary ribs are intercalated. On the outer half of the flank all ribs have

the same thickness and end with uniform marginal nodes. The venter bordered by the slightly surpassing marginal nodes is even and smooth.

**Description** The diameter of the holotype and all other specimens isolated from the Latemar slab do not exceed 40 mm. The specimen from the Ladinian GSSP section at Bagolino (Fig. 3/7a,b, PIMUZ 36534) has a diameter of at least 70 mm. In specimens of the new species with a diameter between 26 and 43 mm the umbilical width varies between 32 and 38%. The number of primaries and respective secondaries per half whorl varies between 8 and 10 and 17.5 and 21, respectively (Table 1). The section of the ribs is semicircular and in the specimens from Latemar the furrows between the ribs are narrow. The ventral smooth interruption of the ribs is narrow in the holotype (Fig. 3/1b) or considerably broader as in specimen PIMUZ 36537 (Fig. 3/4b). The marginal nodes of both sides have generally a corresponding position (not so in specimen PIMUZ 36534 in Fig. 3/7 from Bagolino).

Because of complete recrystallization of the specimens no suture or septa are visible. For this reason it is not possible to decide whether the ammonoid from Latemar is adult or juvenile or if it represents the inner whorls of an adult bigger form.

**Comparison** In *Chieseiceras* (genus established by Brack and Rieber 1986) the size of a specimen is critical for the habit. Especially in *Chieseiceras chiesense* (Mojsisovics 1882) the development of the ribs varies considerably during ontogeny. Therefore, only specimens of about the same size should be compared. *Chieseiceras dolomiticum* is more evolute and has broader ribs than *Ch. chiesense* and *Ch. perticaense*. *Chieseiceras pemphix* (Merian, 1854) is more evolute and has a broader smooth venter. However, this form could be judged only from the illustrations in Hauer (1855) and Mojsisovics (1882) because the type itself is no longer available.

**Discussion** As none of the available specimens or fragments of *Ch. dolomiticum* points towards a diameter exceeding 70 mm and because no suture is visible, the size of an adult specimen remains uncertain.

**Occurrence** From its stratigraphic position *Ch. dolomiticum* is considered a late Anisian form (*N. secedensis* Zone). The specimens from Bagolino were found at the 61.2 m level in a layer above the Tc-tuff interval, i.e., in the middle part of the *N. secedensis* Zone (Fig. 2). The position of the Latemar specimens is less clear but they could possibly be somewhat younger.

**Remarks** The ammonoid species *Ammonites pemphix* Merian, 1854 is here included into the genus *Chieseiceras* Brack and Rieber, 1986. Based on the stratigraphic

succession we propose the following, for ammonoids common, evolutionary trend from more evolute to more involute forms of *Chieseiceras*: *Chieseiceras pemphix* → *Ch. dolomiticum* → *Ch. chiesense* → *Ch. perticaense*.

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